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# Conserved stomatal behaviour under elevated CO<sub>2</sub> and varying water availability in a mature woodland

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### **Summary**

- 1. Rising levels of atmospheric  $CO_2$  concentration ( $C_a$ ) and simultaneous climate change profoundly affect plant physiological performance while challenging our ability to estimate vegetation-atmosphere fluxes. To predict rates of water and carbon exchange between vegetation and the atmosphere, we require a formulation for stomatal conductance ( $g_s$ ) that captures the multidimensional response of stomata to changing environmental conditions. The unified stomatal optimization (USO) theory provides a formulation for  $g_s$  with the ability to predict the response of  $g_s$  to novel environmental conditions such as elevated  $C_a$  (e $C_a$ ), warmer temperatures and/or changing water availability.
- 2. We tested for the effect of  $eC_a$  and seasonally varying climate on stomatal behaviour, as defined by the USO theory, during the first year of free-air  $CO_2$  enrichment in a native eucalypt woodland (the EucFACE experiment). We hypothesized that under  $eC_a$ ,  $g_s$  would decrease and photosynthesis ( $A_{net}$ ) would increase, but fundamental stomatal behaviour described in the USO model would remain unchanged. We also predicted that the USO slope parameter  $g_1$  would increase with temperature and water availability. Over 20 months, we performed quarterly gas exchange campaigns encompassing a wide range of temperatures and water availabilities. We measured  $g_s$ ,  $A_{net}$  and leaf water potential ( $\Psi$ ) at mid-morning, midday and pre-dawn ( $\Psi$  only) under ambient and  $eC_a$  and prevailing climatic conditions, at the tree tops (20 m height).
- 3. We found that  $eC_a$  induced a 20% reduction in stomatal conductance under non-limiting water availability, enhanced mid-morning  $A_{\rm net}$  by 24% in three out of five measurement campaigns and had no significant effect on  $\Psi$ . The parameter  $g_1$  was conserved under  $eC_a$ , weakly increased with temperature and did not respond to increasing water availability.
- 4. Our results suggest that under  $eC_a$  and variable rainfall, mature eucalypt trees exhibit a conservative water-use strategy, but this strategy may be modified by growth temperature. We show that the USO theory successfully predicts coupling of carbon uptake and water loss in future atmospheric conditions in a native woodland and thus could be incorporated into ecosystem-scale and global vegetation models.

**Key-words:** drought, *Eucalyptus tereticornis*, instantaneous transpiration efficiency, photosynthesis, rising CO<sub>2</sub>, stomatal conductance, water-use efficiency

### Introduction

Stomata regulate water and carbon exchange between the vegetation and the atmosphere by their opening and clos-

ing (Sellers *et al.* 1997; Hetherington & Woodward 2003; Buckley & Mott 2013). They provide a means for balancing water loss while allowing access to CO<sub>2</sub> into the substomatal cavity for photosynthesis (Brodribb *et al.* 2009). Stomata respond in a complex and nonlinear way to multiple environmental factors such as temperature (*T*), vapour

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pressure deficit between the leaf and the atmosphere (D),  $CO_2$  concentration in the atmosphere  $(C_a)$  and photosynthetically active radiation (Jarvis 1976: Cowan 1982: Damour et al. 2010). Ecosystem and land surface models require reliable modules to describe stomatal behaviour in response to these climatic variables (Sellers et al. 1997; Kleidon 2004; Damour et al. 2010), particularly under current global change scenarios with rising levels of  $C_a$  (De Kauwe et al. 2013) and simultaneous climate change (IPCC 2014). Despite this, some stomatal models still have difficulty predicting stomatal responses under future global change scenarios (Buckley & Mott 2013; Medlyn et al. 2013; De Kauwe et al. 2014).

Plants adjust their physiological performance to directional changes in environmental conditions to maintain or improve their growth and survival under the novel conditions. While stomata typically close in response to rising  $C_{\rm a}$ , there is variation in the magnitude of their response (Saxe, Ellsworth & Heath 1998; Franks et al. 2013). Stomatal responses to elevated  $C_a$  may involve overall reductions in stomatal conductance as has been frequently observed (Ainsworth & Rogers 2007; Franks et al. 2012; but see Leakey et al. 2006). However, the response to elevated  $C_a$  could involve a change in the magnitude of the intrinsic response of  $g_s$  to a strong environmental stimulus such as D (Heath 1998; Maherali, Johnson & Jackson 2003). Despite understanding the mechanisms of the stomatal closure response (Buckley, Mott & Farquhar 2003), an intrinsic change in stomatal behaviour in elevated  $C_a$ has rarely been specifically assessed (although see Berryman, Eamus & Duff 1994).

Stomatal optimization theory is based on the idea that stomatal opening to allow CO2 uptake inevitably comes at the expense of H<sub>2</sub>O loss (Cowan & Farquhar 1977; Orians & Solbrig 1977; Cowan 1982). Thus, stomata should maximize photosynthetic uptake (A) minus the carbon cost of water used in transpiration ( $\lambda E$ , where  $\lambda$  is the marginal carbon cost of water, in mol CO<sub>2</sub> mol<sup>-1</sup> H<sub>2</sub>O). This optimization theory has been applied many times to successfully predict stomatal responses to environmental conditions, including D (e.g. Arneth et al. 2002; Hari et al. 1986; Thomas, Eamus & Bell 1999; Kleidon 2004; Katul et al. 2010; Vico et al. 2013; Buckley & Schymanski 2014), but it appeared that the theory predicted incorrect responses to C<sub>a</sub> (Lloyd & Farquhar 1994). Following optimization theory, Medlyn et al. (2011) proposed a formulation, called the unified stomatal optimization theory (USO hereafter) that correctly predicts the direction of stomatal response to  $C_a$  and has been shown to fit data from a number of high C<sub>a</sub> experiments (e.g. Barton et al. 2012; De Kauwe et al. 2013; Duursma et al. 2013).

It stands to reason that stomatal behaviour should vary according to the water available to the plant (Cowan 1982; Makela, Berninger & Hari 1996; Manzoni et al. 2011, 2013). Optimization theory predicts that  $\lambda$  should increase with decreasing water availability (Cowan & Farquhar 1977; Zhou et al. 2013). The USO parameter g<sub>1</sub> (proportional to  $1/\sqrt{\lambda}$ ) has been shown to decrease with decreasing water availability in grasses (Manzoni et al. 2011; Zhou et al. 2013) and fast-growing plantation trees (Héroult et al. 2013). Nevertheless, in slow-growing trees from drier origins,  $g_1$  was less responsive to changes in water availability than fast-growing trees from mesic provenances (Héroult et al. 2013). Furthermore, Eamus et al. (2013) found that their proxy for  $\lambda$ , calculated at the whole ecosystem level, increased from the wet to the dry season in an arid savanna. A second prediction from optimization theory is that  $\lambda$  (and  $g_1$ ) shows a weak dependency on temperature (Duursma et al. 2013). This latter prediction emerges from the dependency of  $\lambda$  on the temperature-sensitive non-photorespiratory  $CO_2$  compensation point  $(\Gamma^*,$ Brooks & Farquhar 1985; Bernacchi et al. 2001), and other temperature-sensitive considerations like water viscosity (Prentice et al. 2014; Lin et al. 2015).

While measurements of the variation in  $g_s$  in response to environmental variability in the field can be achieved with seasonal measurements of gas exchange, it is more difficult to test how different Ca levels affect stomatal behaviour in natural conditions. Here, we utilized a free-air CO2 enrichment (FACE) experiment to test the effect of rising  $C_a$  and seasonal climate on stomatal behaviour. We quantified the effect of rising  $C_a$  on  $g_s$  during the first year of exposure to elevated C<sub>a</sub> in a mature (>80 year old), native Eucalyptus woodland under seasonally varying water availability and temperature. Measurements during the first year of the Eucalyptus FACE experiment ('EucFACE') allow testing for the effect of elevated  $C_a$  on  $g_s$  and its coupling with  $A_{net}$ while assessing interactions with a variable climate. The EucFACE experiment is the first of its kind established in a native woodland in a seasonally water-limited climate. This, together with the uniqueness of the facility (Figure S4, Supporting information), provides the opportunity to test whether the USO theory can predict stomatal regulation under rising  $C_a$  in native old tall trees (>20 m). We specifically hypothesized that (i) in mature Eucalyptus tereticornis Sm. trees in natural conditions, elevated  $C_a$ would enhance  $A_{net}$  and reduce  $g_s$ , across a range of climatic conditions, (ii) despite the occurrence of stomatal closure in elevated  $C_a$ , the coupling of  $A_{net}$  and  $g_s$  would not be affected by elevated  $C_{\rm a}$  in these mature trees over an annual time frame, that is the USO parameter  $g_1$  would remain unchanged in response to elevated  $C_a$ , and (iii)  $g_1$ would increase with warmer temperatures and greater water availability.

### Materials and methods

MODEL AND THEORY

We evaluated the unified stomatal optimization model (USO) of Medlyn et al. (2011) for estimating stomatal conductance as a function of environmental and photosynthetic parameters. The USO is based on optimization theory following Cowan & Farquhar (1977). In the USO model,  $g_s$  is approximated as

$$g_s \approx g_0 + 1 \cdot 6 \left( 1 + \frac{g_1}{\sqrt{D}} \right) \frac{A_{\text{net}}}{C_a}$$
 eqn 1

where the term  $g_0$  (in mol m<sup>-2</sup> s<sup>-1</sup>) accounts for cuticular conductance and residual stomatal conductance when photosynthesis  $(A_{\text{net}} \text{ in } \mu \text{mol m}^{-2} \text{ s}^{-1})$  approaches zero, D is the leaf-air vapour pressure deficit (in kPa), and  $g_1$  is the slope parameter (in kPa<sup>0.5</sup>). Fitting the USO model allows to test for differences in  $g_0$  and  $g_1$  between  $C_a$  treatments and among measurement campaigns conducted across seasons.

One further implication of the USO is that the increase in instantaneous transpiration (E) efficiency (ITE) in response to elevated  $C_{\rm a}$  (Saxe, Ellsworth & Heath 1998; Ainsworth & Rogers 2007) should be proportional to the increase in  $C_{\rm a}$  (Barton *et al.* 2012; Duursma *et al.* 2013). ITE is the ratio of  $A_{\rm net}$  to E, which can also be written as (eqn 14 in Medlyn *et al.* 2011):

$$ITE = \frac{A_{\text{net}}}{E} = \frac{C_{\text{a}}P_{\text{a}}}{1 \cdot 6(D + g_1\sqrt{D})}$$
 eqn 2

where  $P_a$  is the atmospheric pressure in kPa. Under common prevailing environmental conditions ( $P_a$  and D), if  $g_1$  remained unchanged from ambient to elevated  $C_a$ , then the ratio of ITE between elevated and ambient  $C_a$  should be proportional to the elevated to ambient  $C_a$  ratio. In our case, we expected the ITE to increase from ambient ( $\approx 390 \, \mu \text{mol CO}_2 \, \text{mol}^{-1}$ ) to  $eC_a$  (ambient  $+ 150 = 540 \, \mu \text{mol CO}_2 \, \text{mol}^{-1}$ ) by a ratio of 540/390 = 1.38.

#### SITE DESCRIPTION AND EXPERIMENTAL DESIGN

The EucFACE experiment is located on an ancient alluvial flood-plain about 5 km from the Hawkesbury River in western Sydney (NSW, Australia) at 30 m a.s.l. elevation (33°37′S, 150°44′E). It is located within a 170-ha remnant of native Cumberland Plain woodland. The vegetation, maintained in this location since the 1880s, retains some old-growth trees and some re-growth vegetation, which has been undisturbed for >80 years. The site mean annual temperature is 17 °C, and mean annual precipitation is 810 mm (Bureau of Meteorology, station 067105 in Richmond, NSW Australia; http://www.bom.gov.au).

At the site, total incoming precipitation (P), air temperature (T) and relative humidity  $(H_{\rm R})$  were monitored on a continuous basis and 15-min total (P) and 1-min averages  $(T \text{ and } H_{\rm R})$  were recorded on data loggers (CR3000, Campbell Scientific Australia, Townsville, Australia). There were six humidity and temperature sensors (HUMICAP  $^{\oplus}$  HMP 155 Vaisala, Vantaa, Finland) on the top of the central tower of each study plot. P was monitored with four automated tipping buckets (Tipping Bucket Rain gauge TB4, Hydrological Services Pty Ltd, Liverpool, NSW, Australia), three on top of the central tower at 23-5 m height and one in an open space 100 m away. Atmospheric water vapour pressure deficit (D) was calculated from temperature and humidity measurements.

The soil at the site is a Clarendon loamy sand, a slightly acid, low-organic C, well-drained soil. The loamy sand surface has >75% sand content, and from ca. 50–300 cm depth, the soil is a sandy clay loam with >30% silt and clay. Groundwater depth is 12 m below the soil surface. The vegetation consists of an open woodland with a density from 600 to 1000 trees ha<sup>-1</sup>, a vegetation area index (including leaves and woody stems)  $\leq 2 \text{ m}^2 \text{ m}^{-2}$  and basal area:  $27.6 \pm 2.7 \text{ m}^2 \text{ ha}^{-1}$ , n = 6 plots (Figure S4). There is a single major canopy-forming tree (*Eucalyptus tereticornis* Sm.) and a minor component of *E. amplifolia* Naudin. Further description of the site and particularly its nutrient status is found in Crous, Ósvaldsson & Ellsworth (2015).

EucFACE consists of six 25-m-diameter circular plots (designated hereafter as 'rings'). Each ring comprises 32 vent pipes in a cylindrical frame extending above the tree canopy to 28 m high

(Figure S4). Each vertical pipe is subtended by a control valve and is connected to a circular plenum (Hendrey et al. 1999) through which a computer-controlled system modulates the release of dilute CO2 into the vegetation according to a proportional-integral-differential type of algorithm (Lewin et al. 2009). Within each ring, there are  $39 \pm 3$  canopy trees (mean  $\pm$  SE, n = 6 rings, max = 48, min = 27), with  $\sim$ 17 dominant and co-dominant canopy-forming trees. The vegetation within three rings was exposed to a  $CO_2$  mole fraction ( $C_a$ ) higher than ambient starting in September 2012, whereas the three ambient rings were identical to the CO<sub>2</sub> treatment rings in infrastructure and instrumentation but received ambient  $C_{\rm a}$ , and are used as control plots. At full operation, the target diurnal  $C_a$  for the rings exposed to elevated  $CO_2$  (henceforth 'e $C_a$ ') was 150  $\mu$ mol  $CO_2$  mol $^{-1}$  above ambient Ca. However, Ca was increased gradually for 6 months to this full operation mode.  $C_{\rm a}$  was increased at a rate of ca. 30  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> month<sup>-1</sup> until mid-summer (February Figure S1).

### LEAF GAS EXCHANGE AND WATER RELATIONS MEASUREMENTS

We performed seven campaigns of leaf gas exchange and water potential measurements from April 2012 to November 2013, inclusive. Five of the campaigns coincided with the period when  $C_a$ was enriched above ambient level according to Figure S1, and four campaigns corresponded with the full treatment of ambient +150 µmol CO<sub>2</sub> mol<sup>-1</sup>. Each campaign was carried out over the course of two sunny days, representative of the season (Table S1). Leaf-level CO2 and H2O exchange measurements were performed with four open-flow portable photosynthesis systems (Li-6400, Li-Cor, Inc., Lincoln, NE, USA). We measured photosynthetic  $CO_2$  assimilation  $(A_{net})$  and stomatal conductance to water  $(g_s)$ under 1800 µmol m<sup>-2</sup> s<sup>-1</sup> photon flux density (provided by the in-built Li-6400 red-blue LED lamp). The C<sub>a</sub> in the cuvette varied between CO2 treatments and among campaigns as follows: cuvette  $C_a$  was set to ~390  $\mu$ mol  $CO_2$  mol<sup>-1</sup> in all rings in April and May 2012 (pre-treatment) and in all ambient rings thereafter. For  $eC_a$ rings,  $C_{\rm a}$  in cuvettes was set to 450  $\mu{\rm mol~mol^{-1}}$  (60  $\mu{\rm mol~CO_2~mol^{-1}}$  above ambient) in October 2012, and 540  $\mu{\rm mol~CO_2~mol^{-1}}$  (150  $\mu{\rm mol~CO_2~mol^{-1}}$  above ambient) in February, May, September and November 2013. Gas exchange measurements were taken when maximum and minimum diurnal  $g_s$  were expected: at mid-morning (9.30-11.10 AEST) and at midday (13·00-15·00), respectively (Barton et al. 2012; Héroult et al. 2013). At each campaign, we measured one single cohort of mature fully expanded leaves. During campaigns when two or more leaf cohorts coexisted, we measured the cohort aged 3-8 months. For the four 2013 campaigns (when the  $eC_a$  treatment was 150  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> above ambient in the eC<sub>a</sub> rings), we followed the leaf cohort that had flushed in October 2012 and hence had developed under the eCa treatment. In each ring, canopy access was achieved with permanent 43-m tall freestanding tower cranes equipped with a 35-m-long jib and working gondola (J4010, Jaso Cranes, Idiazábal, Guipuzkoa, Spain). This access allowed measurements to be conducted right at the treetops at a mean height of  $20.1 \pm 0.2$  m (mean  $\pm$  SE for n = 3-4 trees per n = 3 rings per  $C_a$  treatment, per campaign). Once the elevated  $C_a$ treatment in FACE started, measurements were taken simultaneously in pairs of rings comprising one ambient and one elevated CO<sub>2</sub> ring. We measured two leaves per tree in three (April 2012-May 2013) or four (September and November 2013) dominant or co-dominant E. tereticornis trees per ring. Prior to measurements within each ring (or pair or rings), conditions within the chamber were adjusted to the prevailing environment. Chamber humidity and temperature were maintained as close as possible to prevailing conditions during each measurement period.

In each campaign, except April 2012, we measured leaf water potential ( $\Psi$ ) at pre-dawn ( $\Psi_{pd}$ ) and midday ( $\Psi_{md}$ ) and from February 2013 also at mid-morning ( $\Psi_{mm}$ ). In April 2012, there had recently been standing water owing to heavy rainfall prior to measurements; hence, no  $\Psi$  measurements were taken. We took additional measurements of Ψ in mid-summer (January 2013), when soil water availability reached its minimum within the measurement period (October 2013, only  $\Psi_{pd}$ ), and in the middle of the following summer (January 2014, Fig. 1).  $\Psi_{pd}$  was measured on two random leaves from three random trees in each ring.  $\Psi_{mm}$ and  $\Psi_{md}$  were measured in the same leaves used for gas exchange. Water potentials were measured with a Scholander-type pressure chamber (Model 600l; PMS Instrument Company, Corvallis, OR, USA).

#### DATA ANALYSIS

We fit the USO model (eqn 1) to our measurements of gas exchange in different trees measured at different times of day and campaigns: that is, for each tree, we averaged simultaneous measurements from two leaves. In order to test for significant differences (P < 0.05) in  $g_1$  (eqn 1) between  $C_a$  and among sampling dates, we used two-way analysis of variance conducted on the residuals from the USO model fits. This analysis follows the principle that deviation from the overall USO model fit would represent either Ca or campaign effects analogous to a repeated-measures mixed-model ANOVA. We also tested for the correlation between the ITE in ambient vs.  $eC_a$ . To do so, we fitted a linear regression between mean ITE in pairs of ambient and eCa rings measured simultaneously (i.e. same D and  $P_a$ ). This approach rendered 24 pairs of rings measured simultaneously.

For each campaign during which FACE was operating, we tested for differences between ambient and elevated Ca in gas exchange measurements ( $A_{\text{net}}$ ,  $g_{\text{s}}$  and internal to atmospheric CO<sub>2</sub> mole fraction,  $C_i/C_a$ ) and water potential with one-way ANOVA. For these analyses, we considered measurements from different campaigns at different times of the day from each ring as the experimental unit (n = 3 for each treatment level); that is, we averaged simultaneous measurements across leaves in each tree and then averaged measurements across trees in each ring Significance level was set to P = 0.05 with findings with 0.05 < P < 0.10 considered marginally significant. All analyses were performed in R (R Development Core Team 2014) v 3.1.

### Results

### CLIMATE, PLANT WATER RELATIONS ACROSS SEASONS AND CO2 TREATMENTS

The first two measurement campaigns were preceded by a wet summer ('La Niña' event), with heavy rainfall (>100 mm, Fig. 1) falling about 3 weeks prior to the first set of measurements (Table S1). This period was followed by a relatively dry (with respect to long-term trends) autumn, winter and beginning of the spring, when we took the first round of gas exchange measurements under  $eC_a$ (October 2012, with  $eC_a$  60  $\mu$ mol  $CO_2$   $mol^{-1}$  above ambient). Subsequent campaigns, when  $eC_a$  was 150 µmol CO<sub>2</sub> mol<sup>-1</sup> above ambient, occurred during a wet summer (February 2013), a temperate autumn (May 2013) and

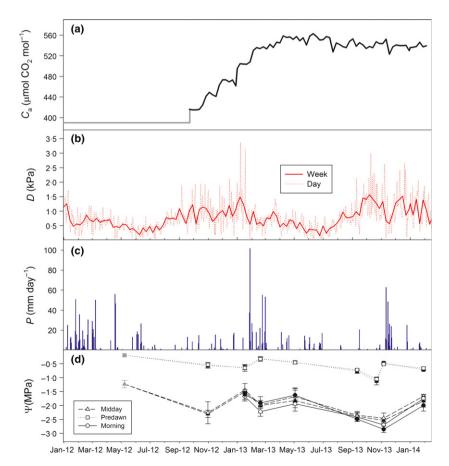


Fig. 1. (a) Step-up increase of  $CO_2$  mole fraction (Ca) in elevated Ca rings (mean weekly  $C_a$ , black line) and approximate ambient Ca during pre-treatment (grey line), (b) mean daily (dotted line) and weekly (continuous line) air water vapour pressure deficit (D), (c) daily precipitation (P) and (d) mean ( $\pm$ SE, n = 3) pre-dawn, morning and midday water potential  $(\Psi)$  in ambient  $C_a$  (open symbols), elevated  $C_a$ (closed symbols) and pre-treatment (grey symbols).

after abnormally warm and dry conditions in both winter and beginning of spring (Fig. 1, September and November 2013, Table S1).

Leaf water potential reflected the rainfall pattern over the year (Fig. 1). The lowest pre-dawn leaf water potential  $(\Psi_{pd})$  was measured at the driest point reached after the longest sustained period without substantial rain (October 2013). A series of four small rain events (10·5 mm of precipitation in total) prior to the following measurement of  $\Psi_{pd}$  (November 2013) was sufficient to boost mean  $\Psi_{pd}$  from  $-1\cdot09\pm0\cdot07$  to  $-0\cdot49\pm0\cdot02$  MPa (mean  $\pm$  SE). Mean  $\Psi_{pd}$ ,  $\Psi_{mm}$  and  $\Psi_{md}$  did not differ between ambient and  $eC_a$  in any campaign (Fig. 1).

### LEAF STOMATAL CONDUCTANCE AND OTHER GAS EXCHANGE PARAMETERS UNDER ELEVATED $\mathcal{C}_{\mathtt{A}}$

Leaf stomatal conductance to water vapour  $(g_s)$  varied seasonally with prevailing climatic conditions (Figs 2 and 3). Maximum  $g_s$  was measured in February 2013 (high water availability and warm temperatures) and minimum in November 2013 (low water availability). In all campaigns,  $g_s$  decreased from mid-morning to midday (Fig. 2) and the magnitude of this drop was largest in spring (in both 2012 and 2013) when  $\Psi_{\rm pd}$  reached its minimum and midday D was 2·5 and 4·4 kPa (in 2012 and 2013, respectively).

We found that an increase of 60  $\mu$ mol  $CO_2$  mol<sup>-1</sup> above ambient (October 2012) did not have a statistically significant effect on  $g_s$ . In February 2013, when  $eC_a$  reached 150  $\mu$ mol  $CO_2$  mol<sup>-1</sup> above ambient, mean  $g_s$  was 20% lower in  $eC_a$  (morning:  $0.43 \pm 0.02$ , midday:  $0.31 \pm 0.05$  mol  $H_2O$  m<sup>-2</sup> s<sup>-1</sup>) than in ambient  $C_a$  (morning:  $0.53 \pm 0.06$ , midday:  $0.43 \pm 0.03$  mol  $H_2O$  m<sup>-2</sup> s<sup>-1</sup>), although this difference was only marginally significant for midday  $g_s$  (F = 5.3, P = 0.084).

Photosynthesis ( $A_{\rm net}$ ) followed a similar seasonal pattern to  $g_{\rm s}$ .  $A_{\rm net}$  peaked in February 2013, when temperatures were warm and water availability was high, and was minimal in May 2012, when it was coldest.  $A_{\rm net}$  always declined with  $g_{\rm s}$  from mid-morning to midday, and this

drop was most pronounced in October 2012 and November 2013. There was significant stimulation of photosynthesis in elevated  $C_a$  (P < 0.05) for three out of five sampling dates when FACE was operating (Fig. 2). First, in October 2012, when  $eC_a$  was 60  $\mu$ mol mol<sup>-1</sup> above ambient,  $A_{\rm net}$  in the morning was  $13.3 \pm 2.7$  and  $19.1 \pm 1.2$  μmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup> in ambient and eC<sub>a</sub>, respectively, although the difference between treatments was not statistically significant (F = 3.8, P = 0.124, Fig. 2). In February 2013, an increase of 150 μmol mol<sup>-1</sup> above ambient  $C_a$  significantly increased  $A_{net}$  in the morning (F = 16.0, P = 0.016). In May 2013,  $A_{\text{net}}$  was not significantly different between Ca treatments. In the following campaign (September 2013), we found that  $A_{net}$  was enhanced under  $eC_a$  both in the morning and at midday (F = 11.4 and 9.9, P = 0.032 and 0.005). Finally, in November 2013, we found significantly higher  $A_{net}$  under  $eC_a$  in the morning (F = 26.1, P < 0.001).

Leaf internal CO<sub>2</sub> mole fraction ( $C_i$ ) in ambient  $C_a$  was conserved across the seasons and was always higher in the morning (285  $\pm$  3  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup>) than at midday (248  $\pm$  5  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup>, n = 7 campaigns). In e $C_a$ ,  $C_i$  was 390  $\pm$  8 (morning) and 340  $\pm$  13  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> (midday, n = 4 campaigns with e $C_a$  = 150  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> above ambient). The ratio of leaf internal to atmospheric CO<sub>2</sub> mole fraction ( $C_i/C_a$ ) did not differ significantly between  $C_a$  treatments, neither in the morning (F = 0·1, F = 0·717) nor at midday (F = 2, F = 0·156) in any campaign.  $C_i/C_a$  was always significantly higher (F = 55·8, F < 0·001) in the morning (0·72  $\pm$  0·01) than at midday (0·63  $\pm$  0·01), when F was higher. There was some seasonal variation in  $C_i/C_a$  (Table S2).

### STOMATAL RESPONSE TO ENVIRONMENTAL PARAMETERS AND COUPLING WITH PHOTOSYNTHESIS

 $A_{\rm net}$  and  $g_{\rm s}$  were broadly correlated, but there was wider scatter in this relationship under  $eC_{\rm a}$  compared to ambient  $C_{\rm a}$  (Fig. 3). Stomatal conductance  $(g_{\rm s})$  was coupled with D (Fig. 3) for all campaigns and  $C_{\rm a}$  levels.  $C_{\rm i}/C_{\rm a}$  decreased with D (Table S2).

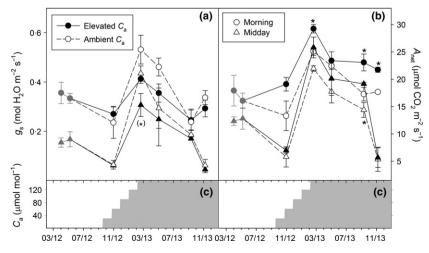


Fig. 2. (a) Mean  $(\pm SE, n = 3)$  stomatal conductance to water  $(g_s)$  and (b) photosynthesis  $(A_{\rm net})$  in the morning (circles) and at midday (triangles) for different  $C_a$  (CO<sub>2</sub> mole fraction) levels: ambient  $C_a$  (open symbols), elevated  $C_a$  (closed symbols) and pre-treatment (grey symbols). Asterisks \* denote significant (P < 0.05) or marginally significant (\*) (0.1 < P < 0.05) differences between ambient and elevated  $C_a$ . (c) Schematic representation of the step-up increase in the elevated  $C_a$  treatment over time.

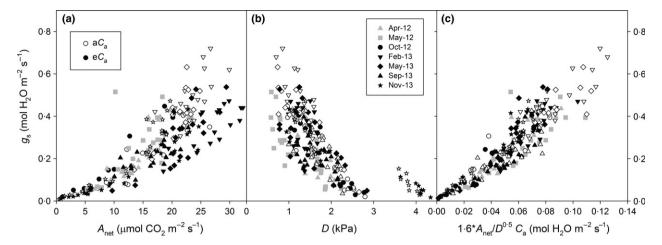


Fig. 3. Relationship between stomatal conductance to water  $(g_s)$  with (a) net photosynthesis  $(A_{net})$  (b) leaf-air vapour pressure deficit (D)and (c) an index based on the USO formulation calculated from D,  $A_{\text{net}}$  and  $CO_2$  mole fraction ( $C_a$ ) for different campaigns and  $C_a$  levels: ambient  $C_a$  (open symbols), elevated  $C_a$  (closed symbols) and pre-treatment (grey symbols). Each point is the mean of two leaves in one tree either in the morning or at midday.

The fitted  $g_0$  parameter, representing the model intercept, was not significantly different from zero for the USO model, in any campaign (Table 1), although the intercept of the overall fit came out as significantly lower than zero, which has no meaningful biological explanation. Given that the  $g_0$  parameter was not part of the theoretical derivation of the USO and that it was not different from zero, it was dropped from subsequent analyses.

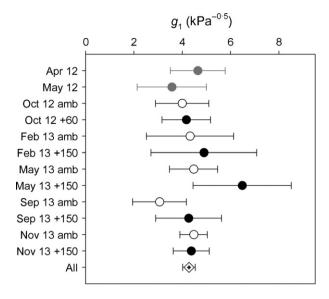
A two-way anova of the residuals of the USO fit showed that there were differences among campaigns in the fitted intercept term (Table 1) and that there were no significant differences in the USO fit neither among campaigns nor between  $C_a$  levels in any campaign (Table 1, Fig. 4). While a slight tendency towards higher  $g_1$  is shown between ambient and eCa (Fig. 4), there was no significant difference in  $g_1$  between  $C_a$  levels in any of the four sampling dates during full elevated  $C_a$  fumigation.

The  $g_1$  parameter (eqn 1) showed some seasonal variation across campaigns (Table 1). Hence, we tested for correlations between  $g_1$  and prevailing climatic conditions of the different campaigns:  $\Psi_{pd}$  and  $\Psi_{md}$ , cumulative precipitation of the 30 days prior to the measurements; minimum and maximum temperatures of the measurement days; monthly mean, maximum and minimum daily temperatures. Among those, we only found that there was a marginally significant positive correlation between  $g_1$  and monthly mean daily temperatures (F = 2.4, P = 0.083,  $R^2$ = 0.19). Contrary to our expectations, we did not find a significant correlation between campaign-based  $g_1$  neither with cumulative precipitation prior to the campaign nor with  $\Psi_{pd}$  (Figure S2).

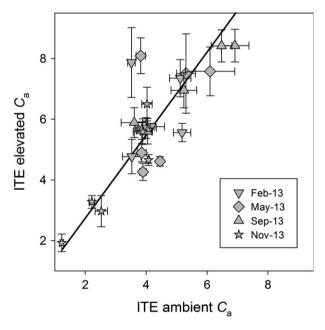
Finally, we tested for the ability of the USO to predict the effect of  $eC_a$  on instantaneous transpiration efficiency (ITE) by fitting a linear correlation between mean ITE in ambient and  $eC_a$  measured simultaneously (same D and  $P_{\rm a}$ ). There was a significant correlation (F = 684.5, P < 0.001, n = 24 pairs of ITE measured under ambient and  $eC_a$  rings simultaneously) with a slope of 1.36. The theoretical value of 1.38 (see Model and theory in the methods section) was well within the 95% confidence interval (CI) of 1.25-1.47 (Fig. 5). We followed a similar approach for the October 2012 campaign (with  $eC_a$ 

Table 1. Results of the two-way ANOVA of the residuals to test for differences between  $C_{\rm a}$  (CO<sub>2</sub> mole fraction) levels, among measurement campaigns and their interaction on the nonlinear fit of the USO formulation. USO model fits (parameter estimates  $\pm$  SE:  $g_0$  in mol H<sub>2</sub>O m<sup>-2</sup> s<sup>-1</sup> and  $g_1$  in kPa<sup>0.5</sup>) for the overall data set ('All') and for each campaign. Parameter values in bold are significantly (P < 0.05) different from 0

Campaign	Effect	Results			Parameters	
		DF	F	P	$g_0$	$g_1$
All	Campaign	1	7.4	<0.001	$-0.038 \pm 0.007$	$4.275 \pm 0.132$
	$C_{\rm a}$	6	0.2	0.672		
	Campaign $\times C_a$	6	1.5	0.186		
Apr-12	Pre-treatment				$-0.057 \pm 0.03$	$4.637\pm0.538$
May-12	Pre-treatment				$-0.017 \pm 0.043$	$3.564 \pm 0.692$
Oct-12	$C_{\rm a}$	1	0.6	0.444	$-0.015 \pm 0.014$	$4.049 \pm 0.334$
Feb-13	$C_{\rm a}$	1	2.2	0.145	$-0.058 \pm 0.047$	$4.841 \pm 0.564$
May-13	$C_{\rm a}$	1	1.6	0.209	$-0.044 \pm 0.032$	$4.69\pm0.456$
Sep-13	$C_{\rm a}$	1	<0.1	0.959	$-0.023 \pm 0.021$	$3.413 \pm 0.406$
Nov-13	$C_{\rm a}$	1	<0.1	0.917	$-0.009 \pm 0.009$	$4{\cdot}242\pm0{\cdot}217$



**Fig. 4.** Fitted parameter with 95% confidence intervals for the unified stomatal optimization model  $(g_1)$  for different campaigns and  $C_a$  (CO<sub>2</sub> mole fraction) levels: ambient  $C_a$  (open symbols), elevated  $C_a$  (closed symbols: 60 and 150  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> above ambient) and pre-treatment (grey). Overall fit for all campaigns and  $C_a$  levels is shown with a different symbol (diamond shape).



**Fig. 5.** Correlation between mean ( $\pm$ SE, n=3–4 trees) instantaneous transpiration efficiency (ITE in μmol CO<sub>2</sub> mmol<sup>-1</sup> H<sub>2</sub>O) from leaf gas exchange in ambient and elevated  $C_a$  rings measured simultaneously in the four campaigns when e $C_a$  was 150 μmol CO<sub>2</sub> mol<sup>-1</sup> above ambient. Line denotes the fitted correlation with a slope of 1·36  $\pm$  0·05 (SE).

60  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> above ambient) a significant correlation (F = 76.5, P < 0.001, n = 5 pairs of rings) with a slope of 1.26 (95% CI: 0.86–1.67 including the expected theoretical value: 450/390 = 1.15). Hence, ITE was correlated proportionally to the amount of CO<sub>2</sub> enhancement.

### **Discussion**

### COUPLING BETWEEN PHOTOSYNTHESIS AND CONDUCTANCE UNDER ELEVATED $\mathcal{C}_{\mathtt{A}}$

We had predicted that stomatal control in ambient and elevated  $C_a$  (e $C_a$ ) would behave according to optimization theory, in particular with the formulation of the unified stomatal optimization model (USO, Medlyn et al. 2011) under a range of conditions, including  $eC_a$ . The USO formulation successfully predicted  $g_s$  under both ambient and  $eC_a$  and parameter  $g_1$  (proportional to the square root of the marginal water cost of carbon gain,  $\sqrt{\lambda}$ ) did not change in response to  $eC_a$ . In agreement with our first hypothesis,  $eC_a$  increased photosynthesis  $(A_{net})$  and marginally decreased stomatal conductance (gs). Elevated Ca profoundly affects physiological performance at different scales, at the leaf scale, and in the short term,  $eC_a$  increases CO<sub>2</sub> supply to the chloroplasts and reduces stomatal conductance. In the longer term,  $eC_a$  could affect leaf and xylem properties related to water supply, but also carbon investments in different organs, which in turn influence photosynthetic function (Domec et al. 2009; Buckley & Schymanski 2014; Tor-ngern et al. 2015), and properties linked to the control of stomata opening (Jones 1998; Katul et al. 2010).

Previously, a possible change in stomatal optimization on forest trees in response to  $eC_a$  had only been evaluated a posteriori, with contradictory findings for different optimization formulations applied to the same experimental data set (Katul et al. 2010; Medlyn et al. 2011; but see Medlyn et al. 2013) and with similar results to ours with a second experimental data set (De Kauwe et al. 2013). Our study is the first that explicitly tests for the effect of  $eC_a$  on optimal stomatal behaviour, in mature trees in natural conditions. Within our study time frame, the USO parameter  $g_1$  was conserved under rising  $C_a$ , as predicted by the USO theory (Medlyn et al. 2011). The USO formulation assumes that stomatal aperture is optimally regulated as if photosynthesis was limited by RuBP regeneration (Medlyn et al. 2013); in our experiment, the  $eC_a$  treatment achieved  $C_i > 350 \mu \text{mol mol}^{-1}$  (except at high D), which is consistent with RuBP-regeneration-limited photosynthesis in Eucalyptus (Lin et al. 2013; Ellsworth et al. 2015). We measured  $A_{net}$  and  $g_s$  along the complete life span of the first leaf cohort developed under  $eC_a$  and yet our results did not show any feedback effects of e $C_a$  on the coupling between  $A_{net}$  and  $g_s$ , resulting from anatomical or physiological changes due to  $eC_a$ (Tor-ngern et al. 2015). Finally, we observed the expected effects of elevated Ca on leaf-level physiology that underlie the theoretical derivation of the original optimization model:  $A_{\text{net}}$  increased while the ratio  $C_i/C_a$  did not change between  $C_a$  levels, providing further evidence of coupling of  $A_{\text{net}}$  and  $g_{\text{s}}$  under different  $C_{\text{a}}$  levels. It remains to be seen whether this persists in the experiment over many years under  $eC_a$ .

One further advantage of the USO theory over previous formulations is that it predicts the theoretical effect of elevated  $C_a$  on instantaneous transpiration efficiency (ITE) from the increase in  $C_a$ . Our analyses showed that the enhancement of ITE under  $eC_a$  agreed with the theoretical value predicted by the USO. These results agree with previous studies where the USO successfully predicted the enhancement of ITE in response to a 60% increase in  $C_a$ in E. saligna (Barton et al. 2012) and cotton plants (Duursma et al. 2013). The ability to predict ITE enhancement under elevated  $C_a$  relies on the assumption that the key USO parameter  $(g_1)$  and the  $C_i/C_a$  do not change from ambient to elevated  $C_a$ , both of which we observed in this study. However, this might not always be the case, at least in conifers (Uddling & Wallin 2012). On the other hand, reduced  $g_s$  in response to elevated  $C_a$  is not uncommon among broadleaved species (Ainsworth & Rogers 2007; but see Keel et al. 2007).

Optimization theory predicts that  $g_s$  should be zero when  $A_{\text{net}}$  is zero (Cowan & Farquhar 1977). Nevertheless, Medlyn et al. (2011) added an intercept term to their formulation, arguing that when  $A_{\text{net}}$  is zero, total conductance to water can be greater than zero due to cuticular transpiration and positive night time g<sub>s</sub> (Barnard & Bauerle 2013). We found that the intercept of the USO was not distinguishable from zero in any campaign. Moreover, when measured under high D (>4 kPa), both A<sub>net</sub> and g<sub>s</sub> neared zero, which suggests that the USO successfully predicts day time  $g_s$  without the need of an additional intercept term.

### OPTIMAL STOMATAL BEHAVIOUR UNDER CHANGING **CLIMATIC CONDITIONS**

The theoretical derivation of the USO model suggests a relationship between  $g_1$  and temperature, through the temperature sensitivity of  $\Gamma^*$  (Medlyn et al. 2011). Assuming the temperature dependence of  $\Gamma^*$  in Eucalyptus reported by Crous et al. (2013), for a 7 °C range in temperature, similar in magnitude to the difference between campaigns with the two most different  $g_1$  values, a  $\approx 22\%$  change in  $\Gamma^*$  can be expected. The  $g_1$  parameter differed by a similar magnitude (24%) between these dates. Across all measurement campaigns, g<sub>1</sub> only weakly co-varied with mean temperature. Under a warmer climate, temperature sensitivity of g1 could have important implications for water and  $CO_2$  fluxes, since small changes in  $g_1$  influence physiological processes such as temperature and D sensitivity of photosynthesis and canopy transpiration (Lin, Medlyn & Ellsworth 2012; Duursma et al. 2014).

The original optimization theory predicts that the marginal water cost of carbon gain decreases with increasing water availability (Cowan 1982; Makela, Berninger & Hari 1996; Manzoni et al. 2013). Analogously and based on the USO theory, we had hypothesized that the  $g_1$  parameter of the USO model would be related to water availability (Zhou et al. 2013). In contrast to these expectations, our results did not show a discernible effect of water availability on  $g_1$ . Our measurements encapsulated a range of water availability spanning from extremely wet conditions when soil water storage was at a maximum across a 4.5-m vertical profile, to a dry period after an unusually warm and dry winter and beginning of the spring when total soil water storage was ca. 25% lower across the same vertical profile (T. E. Gimeno, unpublished data). Nevertheless, we should acknowledge that our study in natural conditions does not allow us to completely disentangle the effects of temperature and water availability on stomatal behaviour. For example, the lowest fitted  $g_1$  corresponded to the coldest campaign (May 2012) and the largest  $g_1$  to the warmest (February 2013), but the latter was also the wettest. In contrast, the driest campaign (September 2013) had an intermediate value for  $g_1$ . Overall, in our study, we did not find a statistically significant trend of  $g_1$  in response to changing water availability, represented by either precipitation of the previous month or pre-dawn water potential. In line with our results, Héroult et al. (2013) and Zhou et al. (2013) showed that in shrubs and trees from subhumid and semiarid origins,  $g_1$  barely changed with increasing water stress. Regarding the effect of future atmospheric conditions on stomatal control, we propose that in mature trees, the marginal water cost of carbon gain will be conserved under increasing  $C_a$ , whereas we speculate, based on our data, that  $g_1$  could increase with warmer mean temperatures.

### **Conclusions**

Stomatal control of simultaneous carbon gain and water loss in Eucalyptus is consistent with the unified stomatal optimization formulation of Medlyn et al. (2011) under seasonally varying climate and more importantly under ambient and future projected Ca. In agreement with the USO theory, the enhancement of ITE under elevated  $C_a$ was proportional to the increase in  $C_a$ . The key USO parameter  $g_1$  (related to the marginal carbon cost per unit water used), in this mature woodland, was conserved under rising  $C_a$ , which allows for predictions of carbon and water exchange under current and future  $C_a$  levels. Also, we did not find evidence for a response of  $g_1$  to changes in water availability, which could suggest that stomatal control of E. tereticornis may have been selected for optimization of C-gain under a drought-prone and extremely variable climate. Nevertheless, this broadly distributed species is likely to be exposed to longer and more severe periods of water scarcity than in this study (Drake et al. 2015). Therefore, we should not generalize this lack of effect of water availability on  $g_1$  to the whole distribution range of this species. Our results suggest that the USO formulation with only a slope term, and no additional intercept, is useful for further incorporation into large-scale models of forest functioning (De Kauwe et al. 2014). Finally, our findings suggest that  $g_1$  could increase with temperature. We venture that this could entail a greater marginal water cost of carbon gain under a warmer climate.

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### Data accessibility

All data used in the manuscript and its supporting information are stored in our institutional archive and publicly available here: http://dx.doi.org/10.4225/35/55b6e313444ff.

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### **Supporting Information**

Additional Supporting information may be found in the online version of this article:

- Fig. S1. Mean (n = 3 rings) hourly CO<sub>2</sub> mole fraction  $(C_a, \text{ black})$
- Fig. S2. Values of the fitted parameter of the USO  $(g_1 \pm SE)$  for each gas-exchange campaign plotted against their corresponding mean ( $\pm$ SE, n = 6 rings) pre-dawn leaf water potential ( $\Psi_{pd}$ ).
- Fig. S3. Mean ( $\pm$ SE, n=3 pairs of rings measured simultaneously) relative effect of the elevated  $C_a$  (CO<sub>2</sub> mole fraction) treatment on photosynthesis  $(A_{net})$  and transpiration in a well-stirred chamber  $(E_s)$  in the morning (red symbols) and in the afternoon (black symbols), with respect to the ambient treatment.
- Fig. S4. (A) overview of the EucFACE site, (B) standing water next to one of the study rings in February 2012, (C) Eucalyptus tereticornis woodland with dry understorey in September 2013.
- Fig. S5. Overview of one of the study rings at the EucFACE experiment during measurements at the tree tops in October 2012.
- Table S1. Summary of the seven gas-exchange campaigns with the exact measurement dates, plots ('rings') measured simultaneously each day, the  $CO_2$  mole fraction ( $C_a$ ) of each ring given as ambient ('Amb'), 60 and 150  $\mu$ mol CO<sub>2</sub> mol<sup>-1</sup> above ambient  $C_a$  (+60 and +150, respectively), the mean maximum and minimum temperature of the exact measurement days (Max T and Min T), the monthly mean maximum and minimum daily temperatures (Daily Max T and Daily Min T) and cumulative precipitation (P) of the previous 30 days.
- **Table S2.** Mean ( $\pm$ SE, n = 6 rings) values of internal to atmospheric CO<sub>2</sub> mole fraction ratio  $(C_i/C_a)$  in the morning and midday for different campaigns.